

## Evidences of three phases of deformation in the Pre-Cambrian rocks of Raisindri Pahar, Singhbhum, Bihar

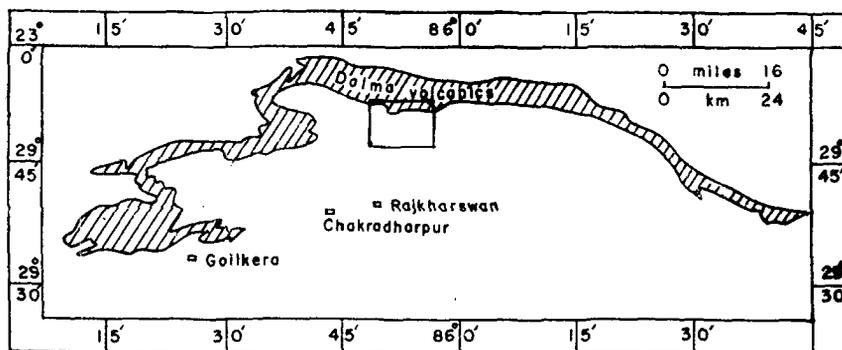
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### Abstract

Earlier workers recognised two phases of deformation. Evidences are presented here which indicate three phases of deformation. The imprint of the last phase of deformation is strongest.

### Introduction and previous work :

The area covers approximately 150 km<sup>2</sup> in the Singhbhum district of Bihar, Eastern India. The rocks are mostly pelitic schists and phyllites with epidiorites belonging to the Iron Ore Series of Dunn (1929). According to Dunn, these rocks underwent two phases of deformation: the first phase corresponds to that of the Dalma syncline; the second phase corresponds to regional folds culminating in an extensive thrusting (the Singhbhum thrust) in which the northern block moved over the southern block. Later, Sarkar and Saha (1962) revised the stratigraphy of this region but essentially retained the above picture given by Dunn. None of these workers, however, carried out any detailed structural analysis of this area.



Location map.

### Structural history :

Detailed structural analysis reveals that there are three distinct phases of deformation. The first phase of deformation produced isoclinal folds on bedding with a perfectly developed axial plane schistosity ( $S_1$ ). Bedding and the first fold features are not preserved on a regional scale. The second phase of deformation involved folding on  $S_1$  along with bedding and was accompanied by the development of a second axial plane schistosity ( $S_2$ ). Structures of the second phase are predominantly preserved in the western and central part of this area. In the northern regions  $S_2$  was not developed, so that the third phase deformation structures were directly superposed on the deformed  $S_1$ -planes.

The third phase deformation features were then superposed on both  $S_1$  and  $S_2$  planes (also on bedding wherever preserved). It caused folding predominantly by

flexural slip. Development of third axial plane schistosity ( $S_3$ ) is scanty except in the shear zones. The third folding movement culminated in shearing along the third axial plane schistosity ( $S_3$ ) within the shear zones in which there is a reversed sense of movement i.e. northern block riding over the southern block.

### Structural features :

The planar structures include bedding ( $S_0$ ) and three consecutive axial plane schistositities ( $S_1$ ,  $S_2$  and  $S_3$ ). Bedding is preserved in only few localities but is important in revealing the nature of the  $S_1$ -planes.  $S_1$ -planes are perfect mineral schistosity planes which both parallel and cross-cut bedding planes and hence are axial plane schistositities parallel to the axial planes of isoclinal folds on bedding. This is confirmed by direct observation, in some of the exposures, of isoclinal folds on bedding with  $S_1$ -planes along their axial planes.

The second generation axial plane schistosity ( $S_2$ ) is most widespread in this area except in the extreme north. This is also a mineral schistosity which has often obliterated the earlier  $S_1$ -planes. The relationship of  $S_2$  to  $S_1$  is very important and subtle. The following points are important in this respect :

(1)  $S_2$ -planes parallel the axial planes of folds on  $S_1$ -planes which may parallel or cross-cut  $S_0$ -planes. Thus, in the same exposures three intersecting S-planes ( $S_0$ ,  $S_1$  and  $S_2$ ) are seen to coexist.

(2)  $S_2$ -planes look coarser than  $S_1$ -planes where the two co-exist.

(3)  $S_2$ -planes are often thoroughly developed to the exclusion of  $S_1$ -planes megascopically. The latter can still be recognised, in a microscope, as relict fold hinges with  $S_2$ -planes as transposition structures.

(4)  $S_1$ -planes show the effect of two later deformations in the form of two axes of folds on this plane whereas  $S_2$ -planes show the effect of a single later deformation. As a result  $S_1$ -planes bear two co-existing crinkle axes and  $S_2$ -planes only one.

Both  $S_1$  and  $S_2$  planes have widely changing orientations throughout the area, whereas the  $S_3$ -planes have approximately constant orientation.  $S_3$ -planes appear as fracture cleavages outside the shear zones and are scanty. Within the shear zone,  $S_3$ -planes are dominant schistosity planes and appear as slip schistosity with a strongly developed striation.

$S_1$  is a mineral schistosity, defined by flaky minerals like muscovite, biotite and chlorite.  $S_1$ -planes are fine and closely spaced.  $S_2$ -planes are also defined by flaky minerals like muscovite and biotite but are less thoroughly developed than  $S_1$ -planes.  $S_3$ -planes are mostly fracture cleavages and transposition schistosity, frequently paralleled by flaky minerals.  $S_3$ -planes are more widely spaced than  $S_1$  and  $S_2$  planes. Evidences of transposition are very well seen in thin sections under the microscope.

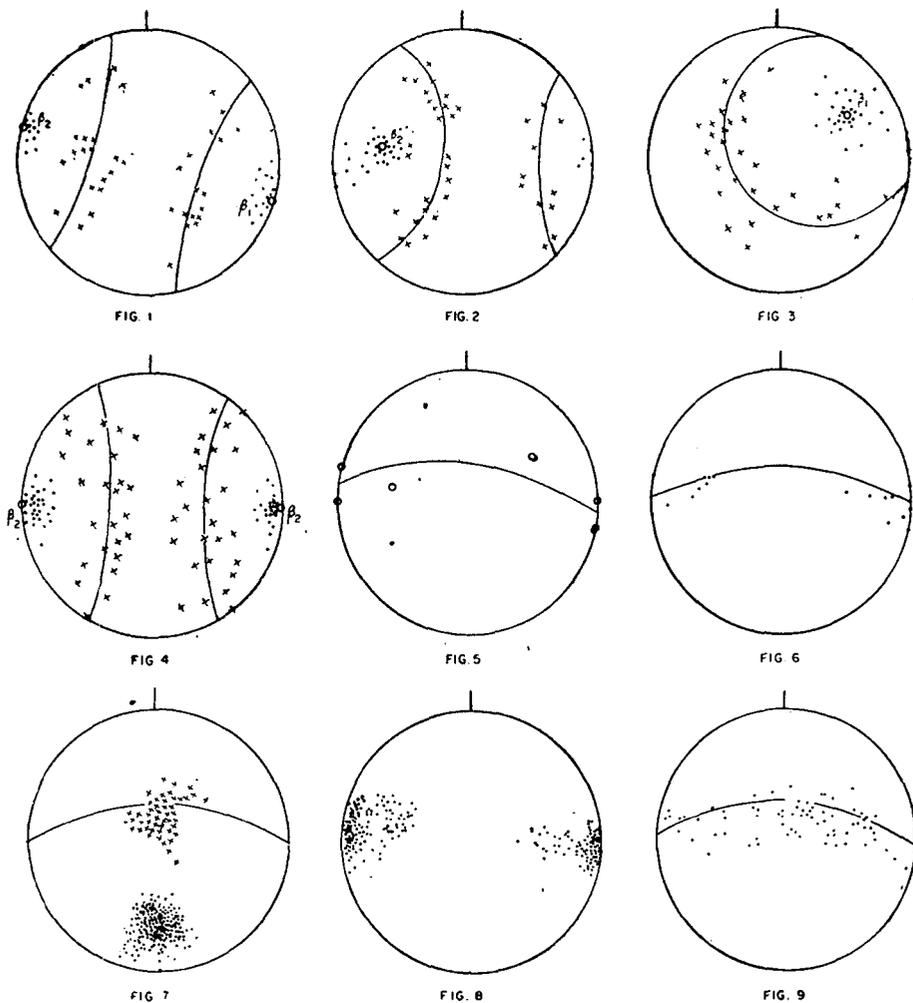
The linear features include axes of crinkles and fold, intersections of planes, mineral lineations and striations. Bedding plane bears commonly three intersection lineations corresponding to the traces of  $S_1$ ,  $S_2$  and  $S_3$  planes. These directions are also commonly paralleled by crinkle axes on bedding.  $S_1$ -planes bear two intersection lineations corresponding to the traces of  $S_2$  and  $S_3$  planes. These are also paralleled by crinkle axes.  $S_2$ -planes bear commonly one intersection lineation corresponding to the traces of  $S_3$ -planes. This is again paralleled by a crinkle axis. At times near or within the shear zones, the  $S_2$ -planes have acted as the shear planes along with the  $S_3$ -planes. Both bear strong striation lines which may be paralleled by mineral lineation and crinkle axis.

Thus, apparently, there are diverse linear features belonging to different genera-

tions of folding movements. Their relative ages were fixed by unrolling in the exposures as also from their patterns on stereograms.

### Folds:

Majority of the mesoscopic and macroscopic folds belong to the third phase of deformation. These are invariably highly asymmetric, overturned to the south. Most of the folds are of flexural slip type except in the thin shear zones. where



Figures 1 to 9.

intense slip folding is common. Both concentric and similar geometry of the folds are present. The overall deformation plan presented by these folds represent southward movement of the northern segments along steeply northerly dipping planes. The regional fold pattern (Fig. 10) exhibited by the Dalma epidiorites, quartzites, phyllites and carbon phyllites in the north are in fact third generation folds with axes plunging towards northeast. This gigantic regional fold is a regional drag fold with a long northern limb, showing a drag sense closely comparable to the drag sense exhibited by the majority of the mesoscopic folds.

### Structural analysis:

For analysis of structural geometry, the area has been subdivided into 4 sub-areas. In most of the subareas,  $S_1$  and  $S_2$  poles show each single great circle distribution; there is a general concentration of poles in the southern parts of the stereograms; this reflects the general overturning of the folds towards south.

The  $\beta$ -poles of such great circles, however, change in orientation. When all the  $\beta$ -poles of  $S_1$  planes (Fig. 5) and of  $S_2$  planes are plotted together it is seen that:

(1)  $\beta_1$ -poles for  $S_1$ -planes change widely in attitude from one subarea to another and lie on a great circle which represents the average orientation of  $S_3$ -planes throughout the whole area.  $S_3$ -planes of the whole area have practically constant orientation (Fig. 7). Thus, all these  $\beta_1$ -poles correspond to the third generation of folds superposed on the second generation of folds on  $S_1$  (Fig. 5).

(2)  $\beta_2$ -poles of  $S_2$ -planes show little scattering and are subhorizontal to low plunging. This is caused by the single folding episode (i.e. third generation folds) affecting the  $S_2$ -planes (Fig. 6).

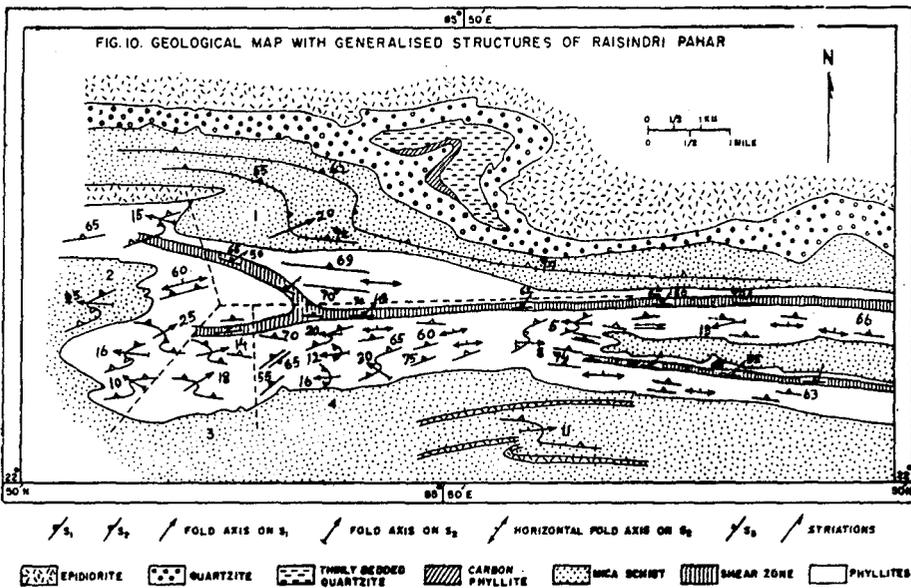


Figure 10.

There are innumerable linear features in these rocks, which present a very complex pattern even in any one subarea. These lineations have been classified according to their ages i.e. successive phases of deformations. Fig. 7 represents a simple picture of striations mostly on  $S_3$ -planes and sometimes on  $S_2$ -planes also. The folds on  $S_2$  are highly asymmetrical with the northerly dipping limb very long and hence majority of the striations are noticed on this limb and are subparallel to the striations on  $S_3$ . Fig. 7 represents all striations observed throughout the whole area and gives a uniform northerly direction of slip which is almost down dip. Fig. 8 represents all  $S_2$ - $S_3$  intersection lineations (which are commonly paralleled by a crinkle axis on  $S_2$ ) throughout the whole area. This also represents a simple picture of subhorizontal to low-plunging ESE-WNW axes of third generation folds on  $S_2$ . Fig. 9 represents traces of  $S_0$  and  $S_1$  planes on  $S_3$  for the whole area. These

lineations are commonly paralleled by axes of crinkles on  $S_0$  and  $S_1$  planes. The pattern shows a scattering along a great circle which represents the average orientation of  $S_3$  planes. Superposition of  $S_3$ -planes on the already folded  $S_1$  and  $S_0$  planes obviously caused this geometrical scattering of the third generation fold axes on  $S_0$  and  $S_1$  planes. Besides these, there are a host of lineations which show interesting patterns:

(a) axes of crinkles on  $S_2$  and  $S_2$ - $S_3$  intersection lineations which are parallel to the  $\beta$ -poles of the subareas and are hence uniform in plunge. These represent the third generation fold axes.

(b)  $S_1$ - $S_2$  intersections which are spread along small circle around the corresponding  $\beta$ -pole of each subarea (Figs. 1-4).

### Regional Synthesis:

The present structural analysis establishes three phases of deformation, hitherto unrecognised in the earlier works. Each phase of deformation is accompanied by the development of characteristic planar and linear features; each of the earlier features are deformed by the later ones. Since the deformation of the third phase is most wide-spread and its features thoroughly preserved, its kinematic pattern could be faithfully unravelled. Within the shear zones, which are products of the third phase of deformation, northerly dipping shear planes ( $S_3$ ) are developed with an almost down-dip slip direction. Outside the shear zones, the third phase deformation structures include drag folds of sizes ranging from microscopic crinkles to regional folds, all overturned towards south. Thus, the overall deformation represents a reversed sense of movement, i.e. the northern hanging block riding over the southern foot block. This picture closely corresponds to the concept of thrusting in the Singhbhum Copper Belt of the earlier workers. Similar structural sequences and kinematic patterns are also obtained in the recent structural investigations of areas to the west and southwest of the present area (Bhattacharyya, 1966; Bhattacharyya and Pasayat, 1971).

### References

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