# Occurrence of liddicoatite-bearing LCT pegmatites in Sirohi region, northwest India and their rare metal potentiality

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Rare and trace element study of pegmatites around Sibagaon area of Sirohi region, Rajasthan, India revealed that the pegmatites are of rare element class and LCT (Li-Cs-Ta) type. Rare metal concentration in these pegmatites is notably high against crustal abundance, with Li ranging from 1007 to 10785 ppm, Rb from 1285 to 9147 ppm, Cs from 36 to 1142 ppm, Ta from 12 to 386 ppm, Sn from 54.22 to 2283.61 ppm and F from 2724 to 48,275 ppm. Integration of EPMA and LA-ICPMS analyses indicated that the main Li-bearing mineral in pegmatites is lepidolite with Li ranging from 21,599 to 28,178 ppm, which is economically worthwhile to process for Li, followed by elbaite and rare mineral liddicoatite, which has not been reported so far in India. Rare earth element (REE) distribution pattern shows enrichment of LREEs than HREEs, indicating that pegmatites are fractionated and thus causing enrichment of rare metals. This reporting of liddicoatite-bearing LCT pegmatites from Sirohi region with high anomalous values of rare metals can therefore be a good prospect for rare metal exploration in India in the present economic scenario. These pegmatites have indicated a syn-collisional signature and may be related to the accretion of Marwar craton with greater Indian land mass along the Phulad Shear zone.

**Keywords:** Liddicoatite, pegmatites, rare metals, trace elements.

THE Aravalli–Delhi Fold Belt in northwestern India is a NE–SW trending, long mountain chain which consists of multiply folded and polymetamorphosed rocks<sup>1–7</sup>. It has always been one of the prime targets for exploration of base metals, radioactive elements, rare earth elements (REEs) and rare metals<sup>8–10</sup>. The closing of Delhi orogeny in northwest (NW) India has led to widespread intrusion of granites and pegmatites. Pegmatites are important sources of rare metals, mainly Li, Cs, Ta, Nb and REEs<sup>11</sup>. The Sirohi region situated in the southwestern sector of NW India hosts widespread granitic and pegmatite bodies. Many workers in the past decades have carried out studies in this region, which were mainly focused on magmatism<sup>12</sup>, granite geochemistry<sup>13</sup>, granite intrusion-

induced deformation<sup>1</sup>, tungsten mineralization<sup>12,14</sup> and distribution pattern of radioactive elements<sup>9</sup>. However, detailed studies on rare metal mineralization have not received much attention.

This study presents the results of systematic geochemical characterization and potentiality of rare metals, mainly lithium, cesium and rubidium of pegmatites in the Sirohi region. Detailed geochemistry of rare metal-bearing mineral phases is also carried out along with first time reporting of the mineral liddicoatite in India. The study also compares the geochemical features and rare metal potentiality of pegmatites of Sirohi region with the world class Tanco pegmatite deposit<sup>15</sup> at Bernic Lake, in southeastern Manitoba, Canada.

The Sirohi region comprises of NE-SW trending ridges formed by Neoproterozoic metasediments of the Sirohi Group<sup>2</sup>. These are fringed on both the eastern and western sides by Erinpura granite and gneisses, and are associated with large-scale faults and shears zones<sup>1,3,16,17</sup> (Figure 1). Around Sibagaon area, located about 15 km towards northeast of Sirohi, a series of pegmatite bodies intruding the rocks of the Sirohi Group are present. Sirohi Group of rocks in the study area is mainly calc-silicate, quartzite, mica schist and phyllite. Pegmatites found in the area are coarse-grained, mainly composed of quartz, feldspar, lepidolite, muscovite and tourmaline (Figure 2). Pegmatites mostly maintain strike continuity with varying width. Marginal areas of pegmatites are medium to finegrained and rich in feldspar, whereas the core is quartzrich. Lepidolite mica is generally associated with feldspar. In most of the pegmatites tourmaline is concentrated in marginal areas, whereas in few pegmatites it is also present in the core part.

The electron probe micro analysis (EPMA) of pegmatite samples on polished thin-sections was carried out (CAMECA SX 100 electron microprobe) in the National Centres of Excellence in Geoscience Researches (NCEGR) laboratory of the Geological Survey of India (GSI) at Faridabad. Operating conditions were 15 keV accelerating voltage and probe current of 20 nA. Analyses were done using a beam diameter of 1 µm, along with all natural standards. The concentration of trace elements in different mineral phases in pegmatite on polished thin-sections was determined using Laser Ablation Inductively Coupled Plasma Mass Spectrometry (LA-ICPMS) at the LA-ICPMS Laboratory, NCEGR, Faridabad, with a CETAC Technologies 213 nm LSX G2 laser ablation unit coupled with a Agilent Technologies 7700x mass spectrometer. The LA-ICPMS was operated at 1350 W plasma power. Helium was used as the carrier gas, with a flow rate of 550 ml min<sup>-1</sup> through the sample cell. Laser ablation conditions were: laser power ~2.5 Mj, with pulse frequencies 10 Hz and spot sizes of 30 µm, carrier gas flow (He + Ar) of  $1.45 \text{ lmin}^{-1}$  in the ICP unit. The nebulizer flow rate was 0.9 l min<sup>-1</sup>. Each analysis consisted of 60 sec of background analyses and 30 sec of ablation

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Figure 1. Geological map of Aravalli–Delhi Fold Belt (modified after Gupta et al.<sup>2</sup> and de Wall et al.<sup>1</sup>).



Figure 2. Field photographs showing (a, b) pegmatites exposed in study area and (c, d) lepidolite in pegmatites.



Figure 3. Photomicrographs of pegmatites showing (a, b) coarse lepidolite (lpd), (c) kink bands in lepidolite, (d) bent twin lamellae of plagioclase feldspar (Plg), (e) perthite (Per) and myrmekite (myr) along the contact of potash feldspar (K-felds) and (f) tourmaline (Trm) being replaced by feldspar.



Figure 4. Geochemical plots of pegmatite showing classification of (a) feldspar, and (b) mica in  $\text{Li}-\text{R}^{2+}-\text{Al}$  diagram of Foster. (c) Si + Li vs Al<sup>vi</sup> + Al<sup>iv</sup> plot of lepidolite showing the substitution mechanism in lepidolite.

time. Data reduction was carried out using the GeoPro software (CETAC). Calibration for the analysis was carried out using NIST 610 and NIST 612 glasses as external standards<sup>18</sup>. Also, <sup>29</sup>Si was used as an internal standard in the analysis. The whole-rock analyses of pegmatite samples was carried out for REE and trace element concentration using Varian 800 ICPMS and Panalytical Axios XRF, at the Chemical Laboratory, Geological Survey of India, Jaipur.

Petrographic studies revealed that the pegmatites are coarse-grained showing equigranular texture with mostly euhedral to subhedral grains and consist of mainly K- feldspar, plagioclase feldspar, quartz, lepidolite mica, muscovite and tourmaline with fluorite as the accessory phase (Figure 3). Plagioclase is corroded at the margins with development of perthitic K-feldspar. At places along the grain boundary of K-feldspar, well-developed myrmekite is presently replacing K-feldspar (Figure 3 e).

EPMA analyses show that plagioclase composition ranges from albite to oligoclase (Figure 4a). For mica, generally major and trace element composition is analysed using EPMA with limitations to detect the lightest elements such as Li, He and H, which are determined using stoichiometry. In the present study, therefore, a

	shorl Elbaite	T7 T8	7.79 37.46	0.10 0.21	7.88 34.58	0.58 0.60	3.18 3.31	0.15 0.08	0.77 4.88	2.53 1.45	1.55 1.82	0.01 0.04	1.31 1.65	0.00 0.00	.814 11.01	2.35 2.39	9.01 99.49			171 10902	2.06 2.69	0.42 1606	4.09 41.42	2.15 17.68	1.45 5.45	6.99 0.09	0.19 186.09	5.95 0.10	0.86 3.53	7.13 154.42	2.09 0.02	5.79 40.62	9.18 79.96	0.96 7.63	4.38 28.68	0.13 3.13	0.29 1.23	0.94 0.27	
	ravite So	T6	6.91 3	0.39	5.38 3	0.14	0.78	0.20	6.75	1.79	1.68	0.02	0.88	0.00	0.86 10	3.07	8.83			3608 2	5.28	4418 16	2.45	9.74 1	3.85	1.34 85	3.32 11	0.88 3	1.64	7.17 1	0.58 7.	0.87	0.90	1.92	8.28	0.60	0.23	0.37	
	Elbaite D	T5	35.17 3	0.16	36.27 3	0.48	2.64	0.12	3.52	0.62	2.23	0.03	0.99	0.01	10.56 1	2.28	95.08 9			72.66	6.24 1	3298 4	106.83 14	0.87 15	6.62 3	0.08	63.76 4	1 10	5.69	136.39 10	1.66	21.71 1	39.13 2	3.47	12.19	0.64	0.24	0.36	
alyses (wt%)	Liddicoatite	T4	35.97	0.13	35.34	0.34	1.86	0.14	0.41	3.06	1.15	0.07	1.46	0.03	10.15	2.09	92.17			3373	7.05	542.11	13.39	616.75	5.39	1035.02	261.82 6 47	0.47 140 50	3.97	90.30	161.10	15.43	27.08	3.39	5.25	2.13	0.99	Ι	
EPMA an	Liddicoatite	T3	37.61	0.15	37.77	0.41	2.23	0.15	0.40	2.94	1.16	0.05	1.50	0.02	10.8	2.49	97.67			6162	0.71	391.69	11.49	4.92	1.46	1.05	0.07	2.06	1.49	33.23	0.14	7.60	10.20	06.0	3.38	0.18	0.09	I	~~~~
	Liddicoatite	T2	38.20	0.12	36.58	0.35	1.93	0.17	0.66	3.14	1.23	0.05	1.48	0.03	10.785	2.57	97.29		es (ppm)	7114	0.82	371.08	4.73	26.13	0.52	0.21	101.78	0.17	2.62	40.22	0.14	8.65	14.13	1.49	4.05	0.17	0.29	I	. 01
	Elbaite	T1	37.09	0.00	37.68	0.42	2.29	0.13	0.68	2.53	1.37	0.06	1.60	0.02	10.76	2.62	97.25	-	wish analys	7572	0.53	258.33	7.55	151.96	0.12	0.01	98.98 164	Ino Ipq	3.67	62.89	I	11.78	18.73	1.71	3.52	0.17	0.11	0.13	
		Point	$SiO_2$	$TiO_2$	$Al_2O_3$	$Fe_2O_3$	FeO	MnO	MgO	CaO	$Na_2O$	$\rm K_2O$	ч	C	$B_2O_3$	$H_2O^*$	Total		LA-IUF	Ľ	Sc	Τi	>	Cr	Ni	Rb	sr v	Zr	qN	Sn	Cs	La	Ce	$\mathbf{Pr}$	рN	Sm	Eu	Gd	Ē
	Muscovite	T6	51.92	0.06	22.55	0.00	0.17	0.03	0.03	0.16	9.87	9.41	0.03	0.05	94.27				31.47		196.57	Ι	172.67	12,509	lbd	lþq	120 2 <i>1</i>	261644	254.84	I	I	14.17	88.00	I	Ι	216.97	I	I	
		Γ1	51.89	0.08	22.42	0.04	0.17	0.15	0.03	0.33	9.78	8.54	0.05	0.03	93.51				8/C,/2 0.72	137.06	0.94	12.90	8.88	13,457	6.11	0.11	1.58	10.801	4026	0.18	0.18	0.01	0.00	I	Ι	0.07	I	Ι	
		L6	52.40	0.04	22.35	0.01	0.17	0.03	0.00	0.22	9.89	8.68	0.01	0.00	93.78				28,178 1.00	199.56		16.95	0.06	14,204	0.94	lbd	0.18	12.9.66	3170	Ι	I	lbd	0.69	I	0.03	0.73	Ι	I	100
wt%)		L5	52.63	0.10	21.56	0.09	0.14	0.09	0.02	0.36	9.65	8.39	0.02	0.02	93.06				990,12 1.73	208 94		I	Ι	13,009	0.32	lbd	0.33	138.55	3431	Ι	lbd	lbd	0.14	Ι	0.09	Ι	;	lbd	
A analyses (	Lepidolite	L4	51.88	0.07	22.67	0.01	0.14	0.07	0.00	0.16	66.6	8.17	0.00	0.00	93.16				606,62 1.43	195.96	0.02	2.59	0.46	13,961	0.24	lþd	0.38	184 00	2301	lpql	0.03	0.05	Ι	Ι	0.11	0.19	lpq	0.06	000
EPM		L3	52.62	0.12	22.50	0.04	0.17	0.10	0.00	0.23	10.00	8.32	0.03	0.01	94.12				20,919 0.26	144.89	0.44	4.89	Ι	12,994	0.18	0.04	- 17050	122.79	2605	0.02	I	0.01	lbd	lbd	Ι	lbdl	I	I	11 1
		L2	51.80	0.00	22.55	0.13	0.10	0.13	0.04	0.36	9.88	8.69	0.04	0.02	93.73			s (ppm)		97 91	1.16	18.45	7.97	130,620	0.31	lþq	164 77	118.53	2880	lpq	0.11	0.26	2.22	2.07	0.73	lbdl	0.21	lbd	
		L1	52.16	0.00	22.69	0.00	0.16	0.13	0.00	0.21	9.74	8.82	0.02	0.00	93.91			MS analyse	20,040 1.67	259.58	0.12	5.01	I	14,428	0.18	lþd	0.30	12.2.19	1580	0.02	lbd	0.02	I	lbdl	0.03	lbdl	lbd	0.08	000
		Point	$SiO_2$	$TiO_2$	$Al_2O_3$	FeO	MnO	MgO	CaO	$Na_2O$	$K_2O$	ц	CI	$P_2O_5$	Total			LA-ICPI	Sc	T: 2	>	Cr	Ņ	Rb	Sr	Y	Zr Nb	Sn	Cs	La	Ce	$\mathbf{P}_{\mathbf{\Gamma}}$	Nd	Sm	Eu	Gd	Tb	Dy	;

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 Table 1.
 (Contd)

			LA-ICP	MS analyses	(mqq)							LA-ICPMS	analyses (ppn	(r			
				Lepidolite				Muscovite		Elbaite	Liddicoatite	Liddicoatite	Liddicoatite	Elbaite	Dravite	Schorl	Elbaite
Tm	I	I	I	lbd	I	I	lbd	lbd	Ho	I	0.02	lbd	0.56	0.01	I	0.12	0.01
Yb	I	I	0.27	0.15	0.12	I	0.12	I	Er	0.01	I	0.02	2.52	lbd	0.08	0.12	0.22
Lu	0.03	Ι	0.09	lbd	0.04	I	I	lbdl	Tm	Ι	I	I	0.12	0.03	Ι	0.13	0.02
Та	20.1	50.72	36.1	39.5	37.3	36.8	39.6	53.8	Чb	0.06	0.00	0.19	1.61	I	I	0.28	0.04
Th	lbd	T	0.01	0.05	0.04	0.14	0.01	lpd	Lu	Ι	I	0.06	I	0.01	lpq	lbd	Ι
N	I	0.17	I	Ι	I	0.20	I	6.78	Та	0.77	0.56	2.1	9.1	0.94	1.7	0.6	6.01
apfu									Th	0.15	0.05	0.63	12.72	0.04	0.04	1.96	1.21
$Si^{4+}$	7.59	7.57	7.61	7.57	7.70	7.63	7.58	7.59	U	Ι	I	0.03	1.96	I	0.04	0.35	0.01
$\mathrm{Ti}^{4+}$	0.00	0.00	0.01	0.01	0.01	0.00	0.01	0.01	apfu								
$AI^{3+}$	3.89	3.89	3.84	3.90	3.72	3.84	3.86	3.88	Si	5.99	6.16	6.05	6.16	6.07	5.79	5.91	5.91
$\mathrm{Fe}^{2+}$	0.00	0.02	0.00	0.00	0.01	0.00	0.01	0.00	Ti	0.00	0.01	0.02	0.02	0.01	0.02	0.05	0.02
$\mathrm{Mn}^{2+}$	0.02	0.01	0.02	0.02	0.02	0.02	0.02	0.02	AI	7.17	6.95	7.16	7.13	7.18	7.04	6.67	6.43
$Ni^{2+}$	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	$\mathrm{Fe}^{3+}$	0.05	0.04	0.05	0.04	0.07	0.06	0.02	0.07
${ m Mg}^{2_+}$	0.03	0.03	0.02	0.02	0.02	0.01	0.03	0.01	$\mathrm{Fe}^{2^+}$	0.31	0.26	0.30	0.27	0.43	0.36	0.10	0.44
$Ca^{2+}$	0.00	0.01	0.00	0.00	00.00	0.00	0.00	0.00	$\mathrm{Mn}^{2^+}$	0.02	0.02	0.02	0.02	0.02	0.02	0.03	0.01
$Ba^{2+}$	0.00	0.00	0.00	0.00	0.00	0.00	00.00	0.00	Mg	0.16	0.16	0.10	0.11	0.18	0.87	1.61	1.15
$Na^+$	0.06	0.10	0.06	0.05	0.10	0.06	0.09	0.04	Ca	0.44	0.54	0.51	0.56	0.44	0.11	0.31	0.25
$\mathbf{K}^{\scriptscriptstyle +}$	1.81	1.84	1.84	1.86	1.80	1.84	1.82	1.84	Na	0.43	0.39	0.36	0.38	0.48	0.71	0.52	0.56
$\mathbf{Al}^{iv}$	0.41	0.43	0.39	0.43	0.30	0.37	0.42	0.41	K	0.01	0.01	0.01	0.02	0.00	0.01	0.00	0.01
$\mathbf{AI}^{vi}$	3.48	3.45	3.44	3.47	3.42	3.47	3.44	3.47	$Fe^{2+}(Y)$	0.31	0.26	0.30	0.27	0.43	0.36	0.10	0.44
									Mg(Y)	0.16	0.16	0.10	0.11	0.18	0.87	1.61	1.15
									Li(Y)	1.06	0.99	0.86	0.50	0.30	1.04	0.50	1.49
									X-vacanc	y 0.12	0.06	0.12	0.04	0.08	0.17	0.17	0.19
bdl, Belo	w detectio	n limit; H <sub>2</sub> C	D*, Calculate	ed from stoic	hiometric (	constraints.											



Table 2. Classification of pegmatites of Sibagaon area, Sirohi district, Rajasthan, India

group Liddicoatite Elbaite Uvite Feruvite Schorl Dravite X-vacant Alkali group group 0.5 0.50 1.00  $(\mathbf{Y})$ Mg<sup>2\*</sup>(Y) Fe "(Y) Mg<sup>b\*</sup>(Y) X-site vacancy Na +(K"

Figure 5. Classification of tourmaline from pegmatites of the study area. *a*, Ternary classification based on dominant occupancy of the X-site. *b*, Ternary classification of alkali group of tourmalines. *c*, Ternary classification of calcic group of tourmalines.



Figure 6. (a) Primitive mantle-normalized multi-element plots and (b) primitive mantle-normalized REE plots for pegmatites of the study area. Normalizing factors from McDonough and  $Sun^{36}$ .

combination of EPMA and LA-ICPMS was used for mica analyses (Table 1). The ternary  $R^{3+}-Li-R^{2+}$  diagram<sup>19</sup> (Figure 4 *b*) shows that most of the micas are lepidolite, with a few muscovite. LA-ICPMS analyses indicate that Li in lepidolite ranges from 21,599 to 28,178 ppm. The main substitution mechanisms which operated in lepidolite is  $2Si + Li \leftrightarrow 3Al^{total}$  indicated by an inverse relation between the Si + Li vs  $Al^{vi} + Al^{iv}$  plot (Figure 4 *c*). This is in agreement with the findings of mica from the Tanco pegmatite deposit<sup>15,20</sup>, which also contains lepidolite showing the same substitution mechanism. For tourmaline analyses also, a combination of EPMA and LA-ICPMS analyses was used. Tourmaline structural formula was calculated using winTcac programme<sup>21</sup>. Table 1 provides the representative analyses of tourmaline. Based on X-site occupancy classification<sup>22</sup>, tourmaline from pegmatites of the study area belongs principally to the alkali group and a few to the calcic group (Figure 5 *a*). The

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(Contd)2.24 3.66 3.25 0.641.1851.1001.771.1771.1583.7482.8333.0690.973.56 0.62 59.13 67.10 46.09 54.87 25.64  $\begin{array}{c} 36.45\\ 76.65\\ 35.43\\ 74.99\\ 68.97\\ 68.97\\ 68.07\\ 50.09\\ 50.09\\ 99.34\\ 91.73\\ 90.73\\ 91.73\\ 167.73\\ 77.60\end{array}$ 111.07 155.79 385.55 385.55 38.52 38.52 38.52 38.56 116.13 116.13 38.56 45.74 79.12 82.35 82.35 82.35 82.35 82.35 88.64 88.64 Та  $\begin{array}{c} 0.02\\$ 0.03 0.05 Lu  $\begin{array}{c} 0.20\\ 0.19\\ 0.16\\ 0.16\\ 0.16\\ 0.16\\ 0.16\\ 0.12\\ 0.10\\ 0.10\\ 0.11\\ 0.01\\ 0.01\\ 0.01\\ 0.01\\ 0.01\\ 0.01\\ 0.00\\$ 0.21 0.33 0.16 0.15 0.140.40 $\begin{array}{c} 0.25 \\ 0.20 \\ 0.27 \\ 0.25 \\ 0.39 \\ 0.12 \end{array}$ 0.36 0.37 0.24 0.36 Yb  $\begin{array}{c} 0.02\\ 0.03\\ 0.02\\ 0.05\\ 0.03\\ 0.03\\ 0.03\\ 0.03\\ 0.03\\ 0.03\\ 0.02\\ 0.01\\ 0.02\\$  $\begin{array}{c} 0.03\\ 0.02\\ 0.02\\ 0.02\\ 0.03\\$ 0.02 0.04 0.02 0.02 0.03 0.03  $\begin{array}{c} 0.03\\ 0.04\\ 0.01\\ 0.05\\ 0.04\end{array}$ 0.03 0.04 0.03 0.03 0.01 Tm Chemical whole-rock analyses of pegmatites showing REE and trace element concentration (ppm) of representative samples 0.10 0.16 0.12 0.170.100.140.150.150.090.130.130.130.150.130.16 0.140.140.140.24 0.06 0.17  $\begin{array}{c} 0.14\\ 0.09\\ 0.20\\ 0.14\\ 0.08\end{array}$ 0.10 0.27 0.10 0.20 0.17 0.13 Ξ 0.140.11 0.27 0.11 0.10 0.19 0.20 0.36 0.17 0.29 0.140.16 0.070.040.060.050.050.030.030.030.030.030.030.030.030.030.030.030.030.030.040.050.060.050.060.030.06 0.05 0.04 0.06 0.040.080.020.040.040.020.020.080.050.050.030.10 0.03 0.04 0.07 0.04 0.06 0.04 0.06 0.06 0.05 0.05 0.06 0.03 0.09 0.05 0.06 0.06 0.09 Но 0.01 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2283.61 176.03 424.29 112.25 576.21 479.58 395.53 241.07 211.88 441.46 367.38 70.91 241.61 296.79 474.87 329.27 310.81 394.28 455.62 420.08 350.98 464.97 751.56 135.58 340.97 190.87 356.92 160.13 129.91 240.98 730.38 345.81 837.45 206.46 2172.47 717.46 603.62 354.97  $\mathbf{Sn}$ 473. 2250 2250 2250 2250 2250 22550 22550 22550 0114 22550 0255 22550 02550 058 058 058 0.98 1.001.85 2.50 2.50 2.50 2.50 2.50 0.71 0.811.23 0.53 1.04 0.99 1.11 2.50 2.50  $\succ$ 8.21 8.61 6.75 6.75 6.75 8.51 6.33 6.56 6.33 9.83 9.83 7.00 21.93 10.79 22.91 8.35 14.53 12.82 7.82 10.98 18.85 0.10 6.73 7.03 15.18  $\begin{array}{c} 5.19\\ 5.15\\ 5.15\\ 6.29\\ 12.50\\ 5.77\\ 5.77\\ 5.77\\ 6.62\\ 6.89\\ 6.89\\ 8.25\\ 8.25\end{array}$ 6.57 8.55 15.58 6.74 11.02 11.03 0.21 3.31 13.09 9.28 g 250.37 180.42 68.20 822.13 53.70 700.36 40.67 145.28 549.26 539.09 68.50 68.50 19.48 10.32 324.69 72.05 68.38 13.49 56.13 29.14 32.08 58.06 07.46 97.25 17.63 15.94 175.46 8.52 10.9455.21 410.84 72.44 190.13 53.82 10.73 26.46 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Table 3.(Contd)

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**Figure 7.** (a) Nb vs Y and (b) Rb vs (Nb + Y) tectonic discrimination diagrams of Pearce *et al.*<sup>27</sup> VAG, Volcanic arc granites; Syn-COLG, Syn-collisional granites; WPG, Within plate granites and ORG, Ocean ridge granites. The x-axis has been modified to accommodate low values of Y and Y + Nb in (a) and (b) respectively.

alkali group of tourmalines shows a varying range of composition and are classified from elbaite to dravite (Figure 5 b)<sup>22</sup>. Depending upon the predominance of Li<sup>+</sup>, Fe<sup>2+</sup> or Mg<sup>2+</sup> on the Y-site, the calcic group of tourmalines are found to be of liddicoatite-type (Figure 5 c). LA-ICPMS analyses of these tourmalines show Li contents ranging from 2,171 to 10,902 ppm.

Classification of pegmatites was carried out on the basis of rare and trace element concentration (Table 2). Pegmatites were classified into rare element (REL) class<sup>23,24</sup>. Based on the family system of petrogenetic classification, these pegmatites were classified as LCT (Li–Cs–Ta) type<sup>23,24</sup>, with enrichment of Li, Rb, Cs, Ta, Sn, F, P, etc. Li concentration ranged from 1,007 to 10,785 ppm, Cs ranges from 36 to 1142.41 ppm, Rb ranges from 1285 ppm to 9147 ppm and Ta ranges from 12 to 386 ppm (Table 3).

Trace element pattern of pegmatites shows relative enrichment of the light (LREEs) and middle (MREEs) rare earth elements over the heavy rare earth elements (HREEs). The REE pattern displays strong negative Eu anomalies (Figure 6*b*) and a distinct enrichment of Cs, Rb, Nb and Ta with depletion in Sr and Ti (Figure 6*a*). Total REE concentrations are typically <100 ppm. These rare element enrichment and depletion patterns in pegmatites of the study area are also consistent with published data for rare-metal pegmatites of the LCT family<sup>23,25,26</sup>. Trace element discrimination diagrams<sup>27</sup> show that majority of the samples plot in the VAG + syn-COLG field of Y–Nb and in syn-COLG field of (Y + Nb)–Rb diagrams (Figure 7).

Detailed mineral chemistry has shown that pegmatites of Sibagaon area, Sirohi region contain abundant lepidolite as the main lithium-bearing mineral, along with liddicoatite and elbaite. Liddicoatite is a gem mineral and is very rare due to substantial concentration of Li and Ca in it<sup>28</sup>. It is found in abundance at the type locality in Madagascar<sup>28-31</sup>, and it is reported for the first time in India. The occurrence of lepidolite with 21,599 to 28,178 ppm Li in these pegmatites is also significant, as lepidolite is economically worthwhile to process for lithium<sup>15,32,33</sup>. Rare and trace element study indicated that pegmatites of Sirohi region are of LCT-type and REL class, with enrichment of rare metals, mainly Li, Cs, Rb and Ta (Table 3), which are of much higher concentration against their crustal abundance that is 20 ppm for Li, 3 ppm for Cs, 2 ppm for Ta and 90 ppm for Rb<sup>34</sup>. Pegmatites of the study area can thus be comparable with the Tanco pegmatite deposit in Canada<sup>15,20</sup> which is LCT-type of pegmatite with lepidolite as Li-bearing mineral and shows similar fractionation trend as well as similar enrichment of rare metals. Pegmatites of the study area therefore can be a potential target for locating important resources of rare metals in India in the present economic scenario. Rare and trace element data also show high concentration of Rb, which may be attributed to abundance of muscovite and K-feldspar. Sr and Ti depletion can be explained by fractionation of plagioclase and Ti-oxide. Negative Eu anomaly along with low concentration of Sr can also be correlated with plagioclase fractionation. Low concentration of U, Th, etc. suggests limited crustal signature. Collisional signature is depicted from trace element discrimination diagrams<sup>27</sup> of the pegmatites (Figure 7). This may be attributed to accretion of the greater India land mass with the Marwar Craton that took place along the terrane boundary Phulad Shear Zone<sup>35</sup> located towards east of Sirohi region.

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